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### Abstract

While global climate modeling algorithms are able to capture the mean behavior of the earth's climate system, sub-scale processes are usually unable to be resolved. In particular, Sea Surface Temperature (SST) anomalies can represent information that is not captured in model climatologies or stochastic events in climate dynamics. Historically, these anomalies have been studied in a Linear Inverse Model (LIM) framework with some success in analyzing monthly, regional anomaly predictability. In this framework, parameters and uncertainties have typically been given point estimates. We examine the decadal predictability of global SST anomalies and discuss the ways in which typical point-estimates are unreliable in this setting. Then we show how using a Bayesian approach can incorporate model assumptions and improve parameter estimates and forecast skill.

# **Sea Surface Temperature Anomalies**



# **Problem Description**

- ► Sea Surface Temperature Anomalies are the discrepancy between physical models of ocean temperatures and measured data.
- Modeling anomalies can produce insight into random effects in the climate system and forces not captured by models.
- ► Statistical models of anomalies can provide information on average behavior as well as inherent uncertainty.
- Better understanding of uncertainty can provide more robust forecasts and policy making decisions.

Data

- ► For methodological concerns, we study CESM Large Ensemble control run data.
- $\blacktriangleright$  Data is gathered at  $1^{\circ}$  resolution for 1800 years for a total of roughly  $10^9$  data points ( 1GB of disk space!). ► In order to feasibly work with that data we have to reduce the amount of data while keeping as much information as possible.

# **Empirical Orthogonal Functions (EOFs)**

- EOF Analysis can be used to study the possible spatial patterns that exist in the data, as well as reduce the total amount of data analyzed.
- Suppose that we have M years of data with N points in space. If our original data, X, has size  $M \times N$  then we can decompose it into temporal patterns, T with size  $M \times D$ , spatial patterns, W with size  $D \times N$  and a matrix representing their relative importance, S with size  $D \times D$ .

$$X = TSW$$

► By only keeping the patterns that are most significant we can identify the important dynamics while reducing the size of the problem and spurious noise present in the data.

#### Leading Spatial Pattern of SST Anomalies





Noticeable patterns include possible ENSO and Kuroshio current related variability.

# **Decadal Predictability of Global Sea Surface Temperature Anomalies Dallas Foster**<sup>\*,1</sup>, N. Urban<sup>2</sup>, and D. Comeau<sup>2</sup>

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0.020 0.017 0.014 -0.011800.0 0.005 0.002 -0.001-0.004 -0.007 -0.010

# **Vector Autoregressive Models**

If we believe that only previous values of the variable are needed in order to forecast the future value then we can use a **Vector Autoregression Model** (VAR).

- Suppose that we expect that the value at time t of our timeseries,  $\mathbf{X}_t$  depends only on the previous p times, then we can form a VAR(p) model  $\mathbf{X}_t = A_1 \cdot \mathbf{X}_{t-1} \cdots + A_p \cdot \mathbf{X}_{t-p} + \epsilon$
- $\blacktriangleright$  A<sub>i</sub> are matrices that give weighted averages between values at different years and different spatial patterns. These
- matrices must be estimated for each application.  $\epsilon \sim N(0, \Sigma)$  is additive noise that captures our uncertainty in predictions. • The simplest VAR model is **Persistence**, the forecast that the next value is the same as the current value  $X_{t+1} = X_t$ .

#### Forecast Skill of 1 year lagged Persistance





# **Linear Inverse Model**

A Linear Inverse Model (LIM) is a mathematical model for the description of processes that vary with random motion. The term in climate dynamics dates from at least the late 1980's and assumes that the dynamics of the process can mostly be expressed by the matrix B and the noise can be captured by the matrix Q and Brownian motion  $d\mathbf{W}_t$ .

$$d\mathbf{X}_t = B\mathbf{X}_t dt + Q d\mathbf{W}_t$$

 $\blacktriangleright$  In order to use a LIM we must estimate the values of the matrices B and Q. In order to do so, it is often convenient to formulate a LIM as a VAR model and use Maximum Likelihood Estimates (MLE). Suppose we want to forecast  $\tau$  years in advance.

$$X_{t+\tau} = e^{B\tau} X_t + \epsilon, \qquad \epsilon \sim N(0, \Sigma)$$

# Maximum Likelihood Estimates

Mathematically we can derive estimates for the parameters B and Q that have maximizes the probability of the data occurring. These are called Maximum Likelihood Estimates (MLEs). These estimates are single values called **point-estimates**.

$$\Lambda = \mathbb{E}[\mathbf{X}_t^2] \qquad \qquad \Lambda_\tau = \mathbb{E}[X_t X_{t+\tau}^T]$$

$$\hat{\Sigma} = \frac{1}{N} \left( \mathbf{X}_{t+\tau} - e^{\hat{B}\tau} \mathbf{X}_t \right) \left( \mathbf{X}_{t+\tau} - e^{\hat{B}\tau} \mathbf{X}_t \right)^T \qquad \hat{Q} - e^{\hat{B}\tau} \hat{Q} e^{\hat{B}^T \tau} = -\frac{1}{N} \left( \mathbf{X}_{t+\tau} - e^{\hat{B}\tau} \mathbf{X}_t \right)^T = -\frac{1}$$

These formulas converge to the actual values of the parameters as the amount of data increases. Their usefulness depends on the amount of data (in time) that we have. 1800 yearly data points may not be enough to ensure good estimates.

#### Forecast Skill of 1 year lagged LIM forecast with MLE Estimates





# **Bayesian Estimation**

 $\Lambda_{\tau} = e^{B\tau} \Lambda$ 

$$\hat{B}\hat{\Sigma} - \hat{\Sigma}\hat{B}^T$$

- MLE estimates tend to underestimate the uncertainty in the system and may not be reliable with the amount of data given. Further, we don't get information about the full distribution of the parameters,  $\theta = [B, Q]$ , or the dynamics. **Bayes' Theorem**:
- $\blacktriangleright$  How do we assign prior probabilities,  $\pi_0(\theta)$  for parameters? We try to enforce properties that we think the matrices should have.
- Uninformative Priors  $\theta \sim N(0, 1)$
- Simple to use and computationally inexpensive.
- ► Does not use knowledge of the system, requires more data to be effective.
- ▶ Minnesota Prior for matrix B
- $\bullet B_{ii} \sim N(0,\lambda), \ B_{ij} \sim N\left(0,\lambda\theta\frac{Q_{ii}}{Q_{ii}}\right), \ \lambda \sim \mathsf{Half-C}(0,1), \ \theta \sim$  $\mathcal{U}(0,1)$
- ► Assumes a random walk nature in the dynamics, which we observe in the data.
- ► Is more computationally expensive than uninformative priors, but only scales linearly with number of parameters.
- ► Can produce sparsity in the matrix, which we also expect.
- ► Horseshoe Prior  $\theta_{ij} \sim N(0, \tau \lambda_{ij}), \quad \tau, \lambda_{ij} \sim \text{Half-C}(0, 1)$ Can produce sparsity in matrices.
- ► Flexible for both matrices *B* and *Q*.
- Large number of hyper-parameters needed to estimate along with matrices. May require more data to be as effective.
- LKJ Prior for matrix Q
- Easy to implement, relatively computationally cheap.
- Naturally designed to sample over correlation matrices.
- ► Not much further structure imposed, like sparsity.

# **Numerical Results on Test Data**

### Comparison of MLE and Bayesian Estimation

Priors for B - Q were chosen and their performance was tested again test data where the values of the parameters were known a-priori. For Bayesian estimates the numerical errors listed below are based on the Maximum A Posteriori (MAP) value of the parameters.

Relative Error in Estimating $B$ Matrix							
N	MLE	unif-unif	unif-horseshoe	horseshoe-unif	horseshoe-horseshoe	minn-unif	minn-horseshoe
50	0.993	1.953	1.971	3.000	7.00	0.574	0.543
100	0.999	1.607	2.371	2.242	2.295	0.298	0.300
1000	1.000	1.415	1.412	0.500	0.503	0.288	0.280
Relative Error in Estimating $Q$ Matrix							
50	0.536	2.157	1.511	3.403	4.405	0.710	0.240
100	0.391	1.214	0.890	1.349	0.711	0.761	0.383
1000	0.098	0.168	0.103	0.158	0.088	0.110	0.075

### Forecast Skill of 1 year lagged LIM forecast with Minnesota-Horshoe Prior Estimates





### **Current and future work**

- Climate Scientists should be skeptical when using MLE based estimates, even when abundant data is available.
- Theoretical convergence may be slow to realize.
- Judging convergence on forecast skill is not a measure for spectra of the parameters.
- ► A Bayesian framework can improve forecasts by incorporating prior beliefs about the structure of estimated parameters.
- Spectra and sparsity can both be better approximated using
- monte carlo techniques.





Figure: Matrix plots for errors of MAP estimates of B matrix using various prior distributions  $\pi_0$  for parameters. These priors can influence the structure of the matrix by helping to promote sparsity.



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